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A dynamic coarsening approach to immiscible multiphase flows in heterogeneous porous media

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1 1 Introduction

2 Numerical modelling is a major tool in applied geoscience for optimizing prospect exploitation at
3 minimal risks and cost while accounting for more and more strict environmental constraints (Suslick
4 and Schiozer, 2004). Efficient and accurate modelling of immiscible flows in heterogeneous porous
5 media is important in many fields of application including oil recovery, waste storage (Nuclear or CO₂),
6 hydrogeology for contaminant remedial purposes, geothermal energy recovery, and other energy-
7 related processes. One of the main challenges related to flow modelling in porous media is the
8 heterogeneity of the rock properties which can vary over several orders of magnitude and many length
9 scales, from pore to reservoir scale. It is well-known that the fine-scale heterogeneity, implying strong
10 localization or channelling of the flow can change field-scale processes, leading to early breakthrough
11 or poor recovery that can hinder the global profitability of an EOR project (Tayari et al., 2018). To
12 capture the influence of the medium heterogeneity, the geological models which describe the reservoir
13 rock properties, are generated at high spatial resolutions. Statistical methods are implemented to
14 propagate the few measured seismic/well/outcrop data to the whole domain, resulting in high
15 uncertainties that must be quantified too. Then the engineer must propagate the uncertainties on the
16 output data of interest, such as hydrocarbons recovery, to recommend the best decisions (Abdul Hamid
17 and Muggeridge, 2018). These models, generally, include several equally probable realizations of the
18 same reservoir. Solving multi-phase flow equations on these high-resolution grids of millions of cells
19 is not computationally efficient. Even with high computing power, it is necessary to run several
20 simulations on different independent realizations, as carrying out a single simulation on a highly
21 detailed geological model is not sufficient. Thus upscaling, the process of propagating the properties
22 from a high-resolution model to a model with less resolution, remains necessary whatever the available
23 computing resources are. The important aim of any upscaling method is to capture the effect of small-
24 scale heterogeneities in an averaged sense, with well-controlled loss of information. While additive rock
25 properties, like porosity, may be upscaled using the direct arithmetic averaging, upscaling the
26 permeability and transmissibility is not straightforward due to the non-linear dependence on the fine-
27 scale properties. The situation is far more complex when considering multiphase flows, due to the strong
28 coupling between pressure and saturation equations that are the basis of the viscous fingering
29 mechanism (Ganjeh-Ghazvini, 2019). Coarse scale equations may be different from their fine-scale
30 counterpart, leading to major changes in the simulation workflow.

31 In single-phase flow, there exist different criteria for the classification of upscaling methods, from
32 analytical averaging methods versus numerical flow-based methods, to local versus extended local and
33 global methods. For example, in local methods, the upscaled permeability in each coarse grid block is
34 solely restricted to the effect of the underlying permeability within the grid block. In extended local
35 upscaling methods, the computation region is extended to include a buffer zone, and in global methods,
36 a fine-scale solution is used to compute the upscaled properties. Extensive reviews may be found in
37 the works of Durlofsky (1991), Renard and Marsily (1997), Wu et al. (2002), Farmer (2002), Mourlanette
38 et al. (2020), and Colecchio et al. (2020). There is no mathematical foundation to analyse the quality
39 and accuracy of the upscaling method, and the closest assessment we can reach is to compare the
40 upscaled solutions with a reference base fine solution (Correia et al., 2018; Darban et al., 2020; Preux,
41 2016; Preux et al., 2016). The challenge in upscaling becomes even more critical in multi-phase
42 processes. The upscaling of absolute permeability, or single-phase upscaling, alone cannot model a
43 multi-phase process and may lead to incorrect oil recovery and water breakthrough times in reservoir
44 simulations. Upscaling of relative permeability, or multiphase upscaling, has been the subject of many
45 articles, but there are practical difficulties in this type of upscaling, no matter what method is used
46 (Barker and Dupouy, 1999; Barker and Thibeau, 1997; Christie, 1996; Darman et al., 2002). The
47 difficulty with the classical methods of multi-phase upscaling, like pseudo-functions, is that they mix
48 averaging with numerical discretization issues (Artus et al., 2004; Noetinger et al., 2004). The net result
49 is an unclear mix of techniques that is over-amplified by the nonlinear character of the underlying
50 equations, typically the so-called pressure equation coupled with the saturation transport equation.
51 These issues are extensively discussed in the mentioned articles.

52 Another challenge in multiphase flow simulations arises from the discretization of the governing
53 equations. A desirable discretization method should be accurate, stable, and locally conservative to
54 respect the physical process. Another important feature is the computational efficiency. Low-order
55 methods such as finite volume methods (FV) are currently widely used for solving these equations.
56 They are stable, mass conservative, and computationally efficient. To improve the accuracy of the
57 approximating solution, especially in the presence of highly heterogeneous anisotropic media, there has
58 been a great interest in high order discontinuous Galerkin (DG) methods over the last two decades. The
59 main idea of DG methods, first introduced by Reed and Hill (1973), consists in approximating the
60 solution using discontinuous polynomials localized in each element and a weakly enforcement of
61 continuity between the elements. It has gone through massive developments leading to different
62 formulations of DG methods. The mathematical aspects of DG methods are detailed in the books of
63 Riviere (2008) and Di Pietro and Ern (2011).

64 DG methods are like finite element methods but with discontinuities in test functions. DG methods can
65 also be viewed like finite volume methods, in terms of element-wise approximating functions. But, in
66 DG methods, the solution is generally approximated by a polynomial of degree greater than one and
67 not by piece-wise constant functions like in finite volume methods. We may say that DG methods
68 incorporate the favourable features of finite volume and finite element methods. They are highly
69 parallelizable, flexible in using variable polynomial degrees in different elements, and locally
70 conservative. These important properties made them popular in many fields of applications including
71 flow and transport in porous media. Many researchers have applied discontinuous Galerkin methods
72 for single and multiphase flow problems of reservoir simulations, from the early work of Riviere et al.
73 (2000), where the DG method was applied to a single-phase problem in porous media, to the works of
74 Riviere and Wheeler (2002), Bastian and Riviere (2004), Ern et al. (2010), Bastian (2014), Cappanera
75 and Riviere (2019), and Fabien et al. (2020) for multiphase flows, to name a few. However, DG methods
76 suffer from one main drawback, which is the increase in the number of degrees of freedom, and
77 ultimately the increase in computational time compared to other numerical methods.

78 Besides adopting higher-order discretization methods, adaptive use of higher resolution grids in some
79 parts of the domain is another approach to improve the accuracy of solutions. The idea of adaptive
80 gridding through local spatial refinement goes back to the 1980s in reservoir simulations (Han et al.,
81 1987; Quandle, 1983; Schmidt and Jacobs, 1988). Conversely, nonuniform coarsening methods can
82 also be adaptively applied to reduce spatial resolution according to specific flow features. The
83 agglomeration approaches have many advantageous especially compared to local mesh refinement
84 thanks to the availability of the fine geological model and have been addressed in several works (Ashjari
85 et al., 2007; Durlinsky et al., 1997; Hauge et al., 2012; Li, 1995).

86 The purpose of this study is to propose an efficient and accurate approach that avoids the challenges of
87 multiphase upscaling and maintains a high order of accuracy in the whole domain. To reach this aim,
88 we have developed a scheme to treat different flow regions separately. Considering a two-phase flow
89 process, like a waterflood in an oil reservoir, there is an interface between the bulk of water and oil.
90 Near the front, the interaction of viscous fingering and heterogeneity of porous medium is important to
91 be captured for industrial applications (Abdul Hamid and Muggeridge, 2018). Like in Hauge et al.
92 (2012), we use adaptive coarsening outside the two-phase flow region to concentrate the calculation
93 efforts within the front area. As a result, the use of any multiphase upscaling technique is avoided. The
94 finite volume method in this high-resolution region gives satisfactory results in a computationally
95 efficient way. Far from the front, in single-phase areas, the grid is coarsened, and a linear DG scheme
96 is used to get more accurate total fluxes. To be able to apply different strategies in different regions, an
97 estimation of the location of the saturation discontinuity is needed. There are different fast approaches
98 for interface modelling. Streamline methods are developed as an alternative to common simulation
99 techniques and work based on the reduction of the main transport problem into a series of 1D problems
100 along streamlines (Datta-Gupta and King, 2007). Another approach to track the evolution of interfaces
101 is the Fast Marching Method (FMM), an approach to solving the Eikonal equation very close to
102 Dijkstra's method (Sethian, 1996). In this study, we present a new technique to estimate the position of

103 the saturation discontinuity. The objective of this method is to approximately track the position of the
 104 front, without solving the fully-coupled pressure and saturation equations on the whole grid.

105 The paper is organized as follows. We detail the formulation and implementation of the proposed
 106 algorithm in the first section and then discuss its advantages and limitations through numerical
 107 examples.

108 2 Driving equations

109 In this paper, we consider an immiscible two-phase flow model, for example, water and oil in a
 110 waterflood problem in a porous medium. When modelling the behaviour of fluids flowing in porous
 111 media, mass conservation and Darcy's law are used to derive the system of equations. Considering the
 112 fluids to be incompressible, the conservation equation for each phase is,

$$\phi \frac{\partial s_\alpha}{\partial t} + \nabla \cdot (\mathbf{v}_\alpha) = q_\alpha, \quad \alpha = o, w \quad \text{Eq. 1}$$

113 where s_α is the phase saturation, \mathbf{v}_α is the phase velocity, q_α is the source term, and ϕ is the rock
 114 porosity. Darcy's law for the phase velocity, in the absence of gravitational and capillary forces, reads,

$$\mathbf{v}_\alpha = -\frac{k k_{r_\alpha}(s_\alpha)}{\mu_\alpha} \nabla p \quad \text{Eq. 2}$$

115 with p the local pressure which is equal to water and oil pressure in our case, k the absolute
 116 permeability, μ_α the viscosity of each phase, and k_{r_α} the relative permeability of each phase, usually
 117 defined as a function of saturation. In the numerical examples, throughout this paper, we use a quadratic
 118 relationship, that reads in terms of reduced water saturation, $s_w^* = \frac{s_w - s_{wi}}{1 - s_{wi} - s_{or}}$: $k_{r_w} = (s_w^*)^2$ and $k_{r_o} =$
 119 $(1 - s_w^*)^2$. k_{r_α}/μ_α is called the phase mobility and is denoted by λ_α . Adding both conservation
 120 equations and considering that $s_w + s_o = 1$ gives,

$$\nabla \cdot \mathbf{v} = q \quad \text{Eq. 3}$$

121 with \mathbf{v} the total velocity or the sum of water and oil velocities, and q the total source term. Writing
 122 Darcy's law for the total velocity gives,

$$\mathbf{v} = -k \lambda_T \nabla p \quad \text{Eq. 4}$$

123 where λ_T is the total mobility, the sum of water and oil mobilities:

$$\lambda_T = \lambda_w + \lambda_o = \frac{k_{r_w}(s_w)}{\mu_w} + \frac{k_{r_o}(s_o)}{\mu_o} \quad \text{Eq. 5}$$

124 A substitution of total Darcy's law into the overall mass-conservation law gives the 'pressure equation':

$$-\nabla \cdot (k \lambda_T(s_w) \nabla p) = q \quad \text{Eq. 6}$$

125 To close the model, boundary conditions are imposed:

$$\begin{cases} p = p_D, & \text{on } \partial\Omega_D \\ \mathbf{v} \cdot \mathbf{n} = 0, & \text{on } \partial\Omega_N \end{cases} \quad \text{Eq. 7}$$

126 which correspond to Dirichlet and homogeneous Neumann boundary conditions on the boundaries of
 127 the reservoir domain. \mathbf{n} is the outward unit normal on these boundaries.

128 Writing the conservation equation of water in terms of total velocity \mathbf{v} and water fractional flow f_w
 129 leads to the ‘saturation equation’:

$$\phi \frac{\partial s_w}{\partial t} + \nabla \cdot (f_w(s_w) \mathbf{v}) = q_w \quad \text{Eq. 8}$$

130 to be solved with the initial and boundary conditions:

$$\begin{cases} s_w = s_w^0, & \text{in } \Omega \\ s_w = s_D, & \text{if } \mathbf{v} \cdot \mathbf{n} < 0 \text{ on } \partial\Omega \end{cases} \quad \text{Eq. 9}$$

131 The system of equations described through Eq. 6 to Eq. 9 is nonlinearly coupled. The coupling arises
 132 via the saturation dependant mobilities in the pressure equation and pressure dependent velocities in the
 133 saturation equation. One common approach for solving this coupled system of equations is the
 134 sequential approach. This approach aims at solving the pressure and saturation separately and
 135 sequentially. The main advantage lies in the reduction of the size of the linear systems to be solved.
 136 Another benefit of this approach is that we can mix different discretization methods in the same system.

137 3 Numerical solution methods

138 3.1 Finite Volume discretization for the pressure equation

139 We assume the reservoir domain Ω , with boundary $\partial\Omega$, is partitioned into cells Ω_i . The standard finite-
 140 volume two-point flux approximation to discretize the pressure equation gives:

$$\sum_j v_{ij}^{n+1} = \sum_j T_{ij}^n (p_i^{n+1} - p_j^{n+1}) = q_i^{n+1}, \quad \forall \Omega_i \in \Omega \quad \text{Eq. 10}$$

141 Where the superscripts n and $n + 1$ represent the time step, $p^{n+1} = \{p_i^{n+1}\}$ is the new cell-wise
 142 constant approximated pressure, q_i^{n+1} is the volume average of the total source term q over the cell i at
 143 timestep $n + 1$, and T_{ij}^n is the face transmissibility given by the distance-weighted harmonic average of
 144 $k\lambda_T^n$ values in the two neighbouring cells Ω_i and Ω_j :

$$T_{ij}^n = |\sigma_{ij}| \left(\frac{d_{i,\sigma}}{k_i \lambda_{T_i}^n} + \frac{d_{j,\sigma}}{k_j \lambda_{T_j}^n} \right)^{-1} \quad \text{Eq. 11}$$

145 $|\sigma_{ij}|$ is the area of the face $\sigma_{ij} = \partial\Omega_i \cap \partial\Omega_j$, $d_{i,\sigma}$ and $d_{j,\sigma}$ denote the distance from respective cell
 146 centres to the centre of the face σ_{ij} . The transmissibility depends on time through its implicit dependence
 147 on the local saturation at time t^n .

148 3.2 Finite Volume discretization for the saturation equation

149 The implicit scheme for solving the saturation equation is,

$$\phi_i |\Omega_i| \frac{s_i^{n+1} - s_i^n}{\Delta t} + \sum_{j \neq i} [f_w(s_{ij}^{n+1}) v_{ij}^{n+1}] = q_i^{n+1} \quad \text{Eq. 12}$$

150 where ϕ_i is the porosity of cell i , $|\Omega_i|$ is the volume of cell i , Δt is the time step, v_{ij}^{n+1} is the total flux
 151 coming from the solution of the pressure equation, and $f_w(s_{ij}^{n+1})$ denotes the fractional-flow function
 152 associated with the face σ_{ij} . $f_w(s_{ij}^{n+1})$ is chosen using the upwind scheme,

$$f_w(s_{ij}^{n+1}) = \begin{cases} f_w(s_i^{n+1}), & \text{if } v_{ij} \geq 0, \\ f_w(s_j^{n+1}), & \text{if } v_{ij} < 0. \end{cases} \quad \text{Eq. 13}$$

153 3.3 Discontinuous Galerkin discretization for the pressure equation:

154 To obtain the DG discretization of the pressure equation, Eq. 6 and Eq. 7, we first re-write the elliptic
 155 pressure equation in a mixed form through the introduction of an auxiliary variable $\mathbf{z} := -\nabla p$ (Frank et
 156 al., 2015),

$$\begin{cases} \mathbf{z} = -\nabla p, & \text{in } \Omega, \\ \nabla \cdot (k\lambda_T \mathbf{z}) = q, & \text{in } \Omega, \\ p = p_D, & \text{on } \partial\Omega_D, \\ \mathbf{z} \cdot \mathbf{n} = 0, & \text{on } \partial\Omega_N. \end{cases} \quad \text{Eq. 14}$$

157 By multiplying the first and second equations by smooth test functions \mathbf{y} and ω , respectively, and
 158 integrating by part on a cell Ω_i we get,

$$\int_{\Omega_i} \mathbf{y} \cdot \mathbf{z} - \int_{\Omega_i} \nabla \cdot \mathbf{y} p + \int_{\partial\Omega_i} \mathbf{y} \cdot \mathbf{n} p = 0, \quad \text{Eq. 15}$$

$$- \int_{\Omega_i} \nabla \omega \cdot (k\lambda_T \mathbf{z}) + \int_{\partial\Omega_i} \omega k\lambda_T \mathbf{z} \cdot \mathbf{n} = \int_{\Omega_i} \omega q. \quad \text{Eq. 16}$$

159 For the numerical resolution of Eq. 15 and Eq. 16, we use FESTUNG¹, the Finite Element Simulation
 160 Toolbox for Unstructured Grids, an open-source MATLAB/GNU Octave toolbox, developed as a
 161 package for discontinuous Galerkin methods. For more details of this DG method and its
 162 implementation, the reader can refer to the series of papers that the authors of this toolbox have
 163 published: Frank et al. (2015), Reuter et al. (2016), Jaust et al. (2018), Reuter et al. (2018), and Reuter
 164 et al. (2020). Let us denote by $\mathbb{P}_P(\Omega_i)$ the space of polynomials of degree at most P on $\Omega_i \in \Omega$ and by
 165 $\mathbb{P}_P(\Omega) = \{\omega_i: \bar{\Omega} \rightarrow \mathbb{R}; \forall \Omega_i \in \Omega, \omega_i|_{\Omega_i} \in \mathbb{P}_P(\Omega_i)\}$ the space of discontinuous polynomials on the
 166 partitioned domain Ω . The DG method is used to solve the following system to obtain $p_i^{n+1} \in P_P(\Omega)$
 167 and $\mathbf{z}_i^{n+1} \in [P_P(\Omega)]^2$, for $\forall \Omega_i \in \Omega$, $\forall \mathbf{y}_i \in [P_P(\Omega)]^2$, and $\forall \omega_i \in P_P(\Omega)$:

$$\int_{\Omega_i} \mathbf{y}_i \cdot \mathbf{z}_i^{n+1} - \int_{\Omega_i} \nabla \cdot \mathbf{y}_i p_i^{n+1} + \sum_j \int_{\sigma_{ij}} \mathbf{y}_i \cdot \mathbf{n}_{ij} \begin{cases} \{p_i^{n+1}\}, & \text{on } \varepsilon_\Omega \\ p_D, & \text{on } \varepsilon_D = 0, \\ p_i^{n+1}, & \text{on } \varepsilon_N \end{cases} \quad \text{Eq. 17}$$

¹ <https://github.com/festung/FESTUNG>

$$\begin{aligned}
& - \int_{\Omega_i} \nabla \omega_i \cdot (k_i \lambda_{T_i}^n \mathbf{z}_i^{n+1}) \\
& + \sum_j \int_{\sigma_{ij}} \omega_i \begin{cases} \{(k_i \lambda_{T_i}^n) \mathbf{z}_i^{n+1}\} \cdot \mathbf{n}_{ij} + \frac{\eta \gamma_{ij}^n}{h_{ij}} \llbracket p_i^{n+1} \rrbracket \cdot \mathbf{n}_{ij}, & \text{on } \varepsilon_\Omega \\ (k_i \lambda_{T_i}^n) \mathbf{z}_i^{n+1} \cdot \mathbf{n}_{ij} + \frac{\eta \gamma_{ij}^n}{h_{ij}} (p_i^{n+1} - p_D), & \text{on } \varepsilon_D \\ 0, & \text{on } \varepsilon_N \end{cases} \\
& = \int_{\Omega_i} \omega_i q_i^{n+1}
\end{aligned} \tag{Eq. 18}$$

168 where η is a penalty coefficient, and h_{ij} is the diameter of the face σ_{ij} . ε_Ω , ε_D , and ε_N denote the set of
169 interior faces, Dirichlet and Neumann boundaries, respectively (Frank et al., 2015). For any interior
170 face σ_{ij} and any function $\omega = \omega(x)$, the jump of ω is defined as,

$$\llbracket \omega \rrbracket_{ij} := \omega_i \mathbf{n}_{ij} + \omega_j \mathbf{n}_{ji} = (\omega_i - \omega_j) \mathbf{n}_{ij} \tag{Eq. 19}$$

171 and the weighted average of ω as:

$$\{\omega\}_{ij} = \omega_i w_{ij} + \omega_j (1 - w_{ij}). \tag{Eq. 20}$$

172 Taking $w_{ij} = 0.5$ leads to the simple arithmetic averaging. Based on Ern et al. (2008), to better take the
173 heterogeneity of the porous medium into account we here take:

$$w_{ij}^n = \frac{\alpha_j^n}{\alpha_i^n + \alpha_j^n} \tag{Eq. 21}$$

174 where $\alpha^n = k \lambda_T^n \cdot \gamma_{ij}^n$ is the harmonic average of α^n values in the two neighbouring cells for the interior
175 faces $\sigma_{ij} \in \varepsilon_\Omega$ (Jamei and Ghafouri, 2016):

$$\gamma_{ij}^n = \frac{2 \times \alpha_i^n \alpha_j^n}{\alpha_i^n + \alpha_j^n} \tag{Eq. 22}$$

176 This value reduces to α_i^n in Dirichlet boundaries. To reconstruct conservative fluxes from the DG
177 solution we compute:

$$\mathbf{v}_{ij}^{n+1} = \begin{cases} \int_{\sigma_{ij}} \{(k_i \lambda_{T_i}^n) \mathbf{z}_i^{n+1}\} \cdot \mathbf{n}_{ij} + \frac{\eta \gamma_{ij}^n}{h_{ij}} \llbracket p_i^{n+1} \rrbracket \cdot \mathbf{n}_{ij}, & \text{on } \varepsilon_\Omega, \\ \int_{\sigma_{ij}} (k_i \lambda_{T_i}^n) \mathbf{z}_i^{n+1} \cdot \mathbf{n}_{ij} + \frac{\eta \gamma_{ij}^n}{h_{ij}} (p_i^{n+1} - p_D), & \text{on } \varepsilon_D. \end{cases} \tag{Eq. 23}$$

178 which corresponds to Eq. 18 with the test function $\omega = 1$.

179 3.4 Fast front tracking technique

180 We have developed a fast front tracking technique (FFrT) to estimate the position of the saturation
181 discontinuity without solving the transport equation in each cell of the domain. Considering a two-phase
182 flow problem like a waterflood, we suppose that the saturation takes only two values corresponding to

183 the saturation ahead of the front (typically s_{w_i}) and behind the front (typically s_f), the saturation
 184 corresponding to the Buckley-Leverett shock front. The mobility ratio between both fluids corresponds
 185 to the mobility ratio at the front (Hagoort, 1974; King and Dunayevski, 1989; Noetinger et al., 2004;
 186 Spesivtsev and Teodorovich, 2007; Teodorovich et al., 2011). The idea is that the velocity of the front
 187 is mainly controlled by the interplay between the heterogeneity of absolute permeability and the
 188 mobility jump at the front, and not by the rarefaction wave behind the front that is continuously
 189 spreading during the time. This method is an extreme simplification of the initial problem to compute
 190 a fast estimation of the location of the water-oil interface based on some concepts from pore network
 191 modelling and Buckley Leverett equation. The FFrT method is an initial value problem, meaning that
 192 it computes the position of the interface from the given initial position of the front or the saturation
 193 discontinuity. According to the method of characteristics applied to the Buckley-Leverett equation, the
 194 velocity of the slice of the porous medium with a saturation s , $\mathbf{v}|_s$, is:

$$\mathbf{v}|_s = \frac{\mathbf{v}}{\phi} \frac{\partial f_w}{\partial s} \Big|_s \quad \text{Eq. 24}$$

195 In other words, if at the point \mathbf{x}_f and at time t_1 the saturation is equal to s_f , it is still equal to this value
 196 at the position $\mathbf{x}_f + \Delta t \frac{\mathbf{v}}{\phi} \frac{\partial f_w}{\partial s} \Big|_{s=s_f}$ and at time $t_1 + \Delta t$.

197 3.4.1 Formulation

198 We define ζ_F^n as the set of all faces σ_{ij} where the front crosses the segment $(\mathbf{x}_i, \mathbf{x}_j)$. \mathbf{x}_i and \mathbf{x}_j represent
 199 the cell centres of the cells i and j . \mathbf{x}_j represents the downstream cell depending on the direction of the
 200 flow (see Figure 1). ζ_F^0 is initialized depending on the initial conditions of the transport problem. Now,
 201 consider that we are at time step n of the simulation. After solving the pressure equation to obtain new
 202 pressure and fluxes, using the DG scheme (Eq. 17 - Eq. 18), we advance the front, in the following
 203 detailed steps:

204 (1) The Welge tangent method is used to construct the saturation at the front, s_f , analytically
 205 (Welge, 1952):

$$\frac{\partial f_w}{\partial s} \Big|_{s=s_f} = \frac{f_w(s_f) - f_w(s_{w_i})}{s_f - s_{w_i}} \quad \text{Eq. 25}$$

206 For the case of quadratic relative permeabilities, the equation above gives the saturation at the front as
 207 a function of the viscosity ratio, $M = \mu_o/\mu_w$, as follows:

$$s_f = 1/\sqrt{1 + M} \quad \text{Eq. 26}$$

208 (2) We compute the velocity at the front, $v_{f,i,j}^{n+1}$, using the discretized form of Eq. 24 and the total
 209 fluxes v_{ij}^{n+1} :

$$v_{f,i,j}^{n+1} = \frac{v_{ij}^{n+1}}{|\sigma_{ij}|\phi} \frac{\partial f_w}{\partial s} \Big|_{s_f} \quad \text{Eq. 27}$$

210 (3) Considering that only the normal component of velocity accounts for the local velocity at the
 211 front, the discretized front is advanced according to the ordinary interface evolution equation:

$$\frac{\mathbf{x}_{f,i,j}^{n+1} - \mathbf{x}_{f,i,j}^n}{\delta t^{n+1}} = v_{f,i,j}^{n+1} \mathbf{n}_{ij} \quad \text{Eq. 28}$$

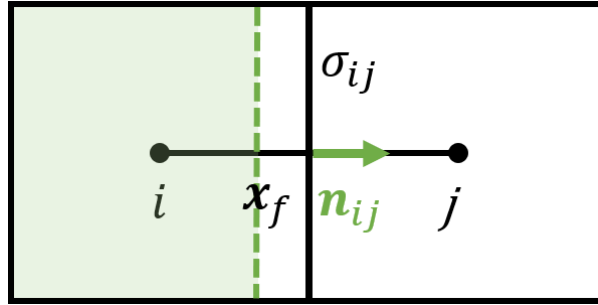
212 δt^{n+1} denotes the sub-timestep for the update of the front position.

213 (4) We define the sub-timestep by posing a condition that measures the smallest time to reach the
214 cell centre:

$$\delta t^{n+1} = \min_{\sigma_{ij} \in \zeta_F^n} \left(\frac{\|\mathbf{x}_j - \mathbf{x}_{f,i,j}^n\|}{v_{f,i,j}^{n+1}} \right) \quad \text{Eq. 29}$$

215 $\mathbf{x}_{f,i,j}^n$ represents the current position of the front between two adjacent cells i and j (see Figure 1)

216 We repeat the steps (3) and 0 until the sum of all δt^{n+1} will be equal to the time step in the main scheme,
217 Δt^{n+1} . Doing so, all the blocks where the front moves to are marked to remain at a higher resolution.



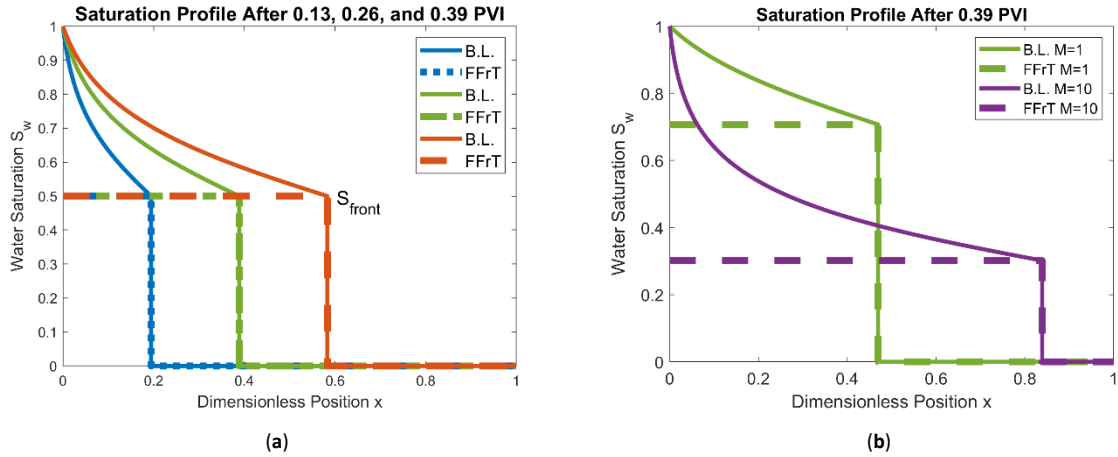
218
219 Figure 1 the two cells and the front position used to define the algorithm

220 3.4.2 Test case validation

221 To verify the proposed FFrT method, we consider two-dimensional water-flood problems, where oil is
222 displaced by the water of the same or different viscosity. Both fluids are incompressible. The reservoir
223 is initially filled with oil. Water is injected from the left boundary of the domain.

224 The Buckley Leverett problem:

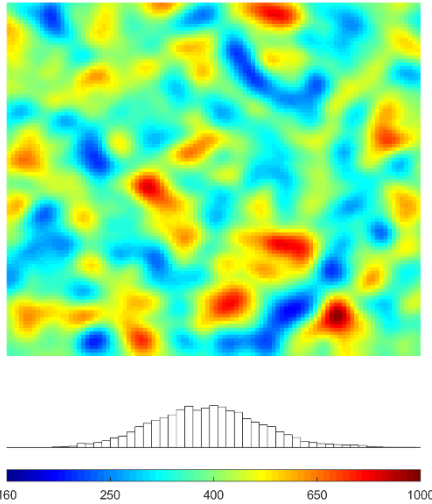
225 We create a simple example to compare the proposed fast front tracking method with the Buckley
226 Leverett (BL) solution. We choose a $10 \times 2 \text{ m}^2$ domain, discretised by a 200×50 Cartesian grid. The
227 permeability and porosity are spatially homogeneous and equal to 0.01 Darcy and 1.0, respectively.
228 Water is injected from the left boundary with a constant rate of $10^{-6} \text{ m}^3/\text{s}$ ($4.3e - 3$ PV/day). The
229 viscosity of oil is set to 10 cp . The viscosity of water varies between three different values of
230 1, 3.33, and 10 cp , in different cases. Figure 2 shows some comparisons of the FFrT method with the
231 BL solution for different viscosity ratio cases and at different time steps.



232 Figure 2 The comparison of the proposed FFrT method with the analytical solution of the BL equation
 233 for (a) a viscosity ratio of $M = \mu_o/\mu_w = 3$, at three different times and (b) for two different viscosity
 234 ratio $M=1$ & 10 , after 0.39 pore volume injected.

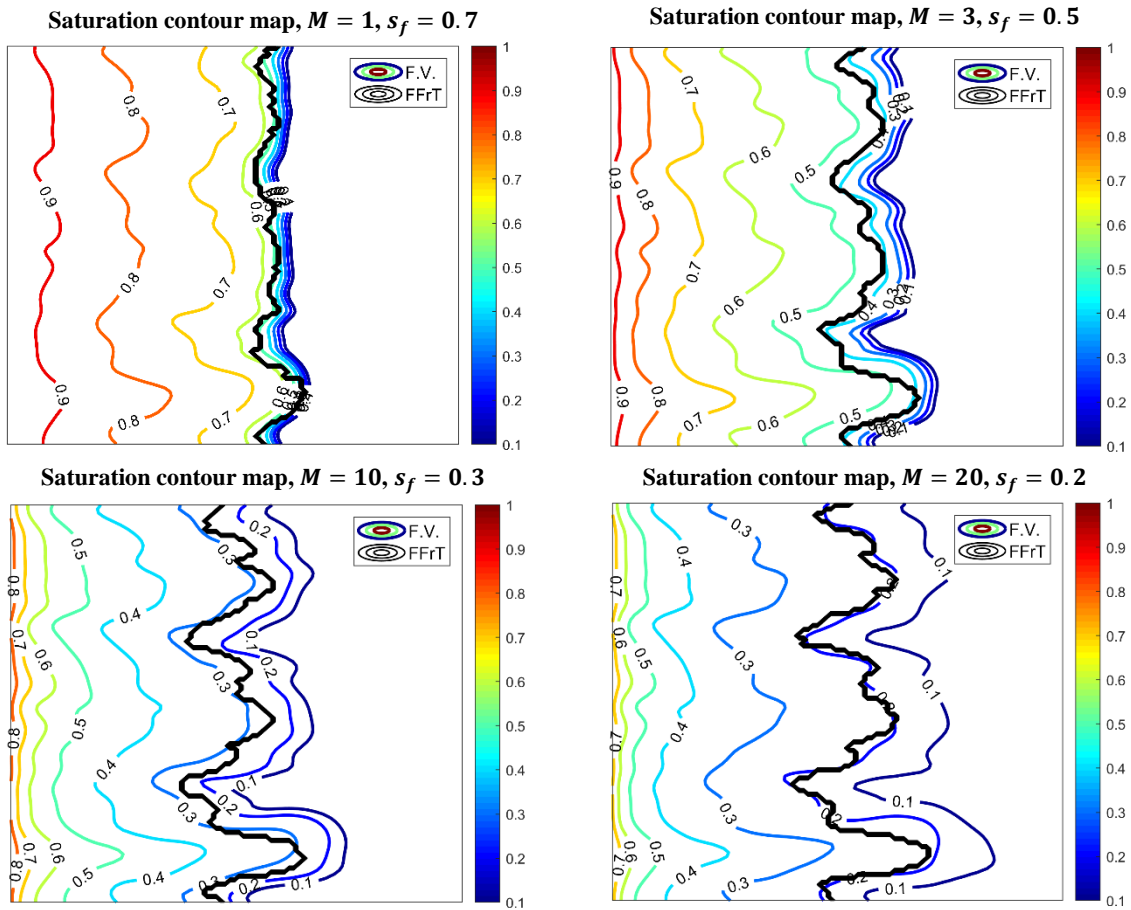
235 **Random generated isotropic medium:**

236 To increase the complexity of the validation cases, we test the method with a smoothly heterogeneous
 237 random permeability field generated with a lognormal distribution, a correlation length of 0.1 , and a
 238 Dykstra-Parsons coefficient of 0.2 , shown in Figure 3. We set Dirichlet conditions on the left and right
 239 boundaries, and a no-flow condition on the top and bottom boundaries. We consider the previously
 240 described two-phase flow problem. We set the viscosity of oil to 10 cp and consider different viscosity
 241 ratios. The original computational grid contains 100×100 cells. Figure 4 shows water saturation
 242 contour maps computed using a standard FV IMPES (implicit pressure - explicit saturation) scheme
 243 and the FFrT method for four viscosity ratios. The analytical saturation of the front is also shown for
 244 each case of the viscosity ratio M , computed using Eq. 26. In the case of a unit viscosity ratio with a
 245 relatively sharp front, the prediction of the FFrT method is in good accordance with the saturation
 246 profile computed using the standard FV scheme. When the viscosity ratio M increases and the front
 247 becomes more distorted, a small difference between the predicted front position and the FV solution
 248 can be observed. Due to numerical errors inherent to both methods and lack of an exact solution it may
 249 be difficult to interpret the difference. It is worth noting that the proposed fast front tracking method
 250 cannot predict the rarefaction wave behind the front and the numerical diffusion of the front. The
 251 essential fact is that the frontal zone is captured correctly. This allows to set-up correctly the adaptive
 252 coarsening strategy with the FFrT method as a criterion, which is our main goal in the proposed
 253 workflow.



254

255 Figure 3 Isotropic randomly generated permeability model with a dimensionless correlation length of
 256 0.1 and Dikstra-Parson coefficient of 0.2 in the logarithmic scale

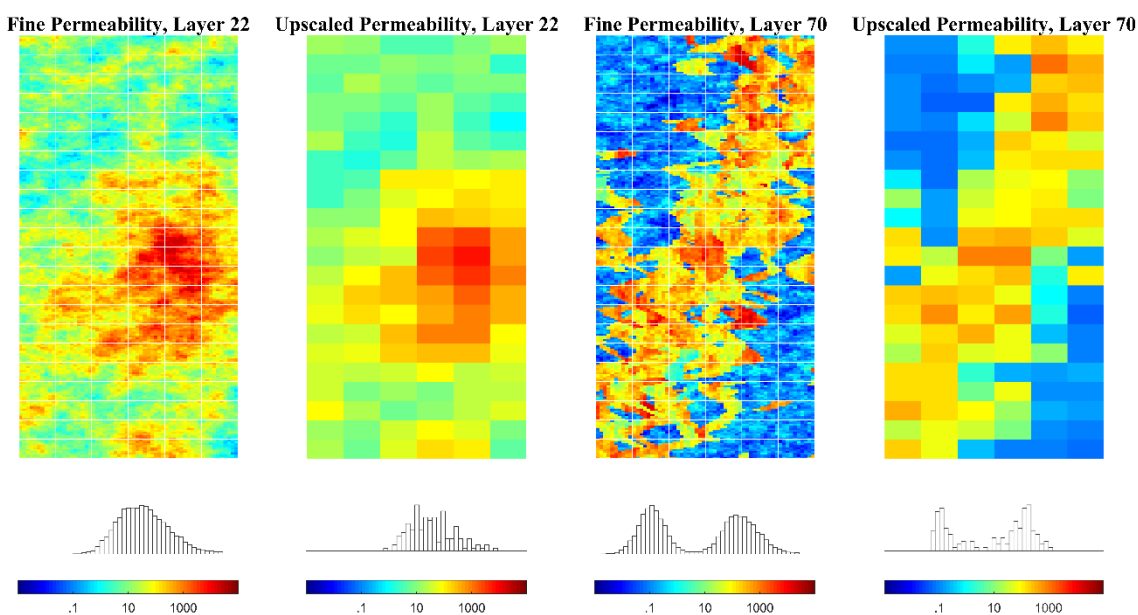


257 Figure 4 Water saturation contour maps for the randomly generated permeability domain shown in the
 258 previous figure with different cases of viscosity ratio, computed by the proposed FFrT method and a
 259 standard FV scheme.

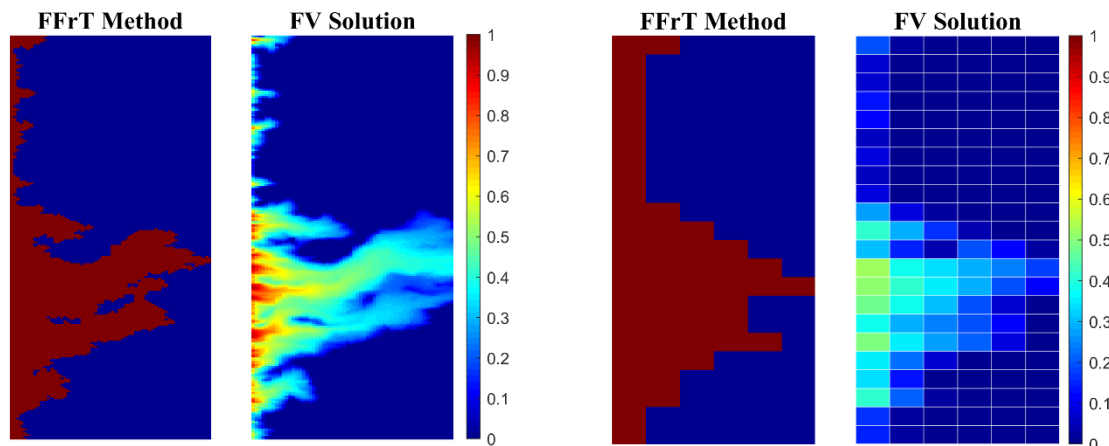
260 **SPE 10 benchmark test:**

261 We consider a two-dimensional Cartesian model with permeability values taken from the second SPE10
 262 benchmark test (Christie and Blunt, 2001). This model contains $60 \times 220 \times 85$ cells, in which the top
 263 35 layers represent Tarbert formation and the bottom 50 layers represent Upper Ness formation. We
 264 consider the same immiscible two-phase flow (water-oil) problem with the same initial and boundary

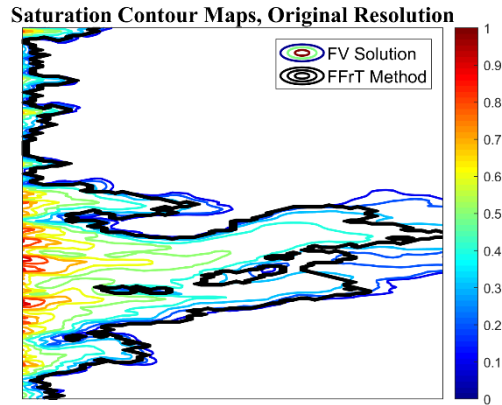
265 conditions as in the previous example, on both the original fine and coarsened grids. We generate the
 266 coarse grid via a uniform agglomeration of the base fine grid, with an agglomeration ratio of 10. We
 267 use a flow-based upscaling method to compute the upscaled permeabilities (Chen et al., 2003), where a
 268 set of representative boundary conditions are imposed at the coarse grid blocks to solve the flow
 269 equation and use the fine-scale pressure and fluxes to compute the upscaled permeabilities. The fine-
 270 scale and upscaled permeability maps for the layers 22 and 70 of the SPE10 dataset are shown in Figure
 271 5. In Figure 6 we compare the FFrT method with a standard finite volume scheme at two resolutions,
 272 for a waterflood problem with a viscosity ratio $M = 10$ in layer 22 of SPE 10 model. At the coarse
 273 scale, the results of FFrT method are very close to the FV solution of the same resolution. In the fine-
 274 scale simulation, the FFrT method predicts the flow patterns very well, and the main difference with
 275 the fine-scale simulation is close to the right boundary, where the front becomes diffusive. For a more
 276 detailed comparison, the superimposed contour maps for the original resolution case are shown in
 277 Figure 7. This figure shows that the predicted front position is very close to the saturation contour line
 278 of 0.3, computed using the FV scheme on the original grid.



279
 280 Figure 5 Permeability maps of the layers 22 and 70 of SPE10 model in the logarithmic scale, before
 281 and after upscaling



282 Figure 6 Water saturation maps for layer 22 of SPE 10 model for a waterflood problem with the viscosity
 283 ratio of 10 at the original and a coarsened resolution, computed by the FFrT and FV methods.



284
 285 Figure 7 Water saturation contour maps for layer 22 of SPE 10 model, for the same time step as the
 286 previous figure, at the original fine resolution.

287 **4 The solution flow chart**

288 We now detail the complete solution procedure at each time step using the previous schemes and the
 289 FFrT technique:

- 290 (1) The pressure equation is solved using the DG scheme with a linear approximation on the base
 291 coarse grid over the whole domain.
- 292 (2) DG conservative fluxes are reconstructed and then the front is moved using the FFrT method.
 293 This step includes solving Eq. 28 in some parts of the domain. The position of the front is then
 294 used to (1) partition the domain into single- and multiphase-phase flow regions and (2) use a
 295 higher resolution grid, equal to the resolution of the original fine grid or an intermediate
 296 resolution, where the front moves. One immediate advantage of domain partitioning is to solve
 297 the saturation equation only in the multi-phase flow region. As a result, no multiphase upscaling
 298 method is needed. We use the dynamic coarsening approach of Hauge et al. (2012) to use a
 299 higher-resolution grid where the criterion is met.
- 300 (3) Flow and transport problems are solved in the multiphase flow part of the domain, on an
 301 adaptively coarsened grid generated at the previous step. Depending on the type of faces,
 302 boundary or interior ones between single-phase and two-phase flow regions, the boundary
 303 pressure or the DG conservative fluxes are set as boundary conditions. These boundary
 304 conditions are used to solve the pressure and transport problems using the FV method. Except
 305 for the DG and the FFrT, our methodology use tools for the upscaling, the dynamic mesh
 306 coarsening, and the FV resolutions that are available in the MATLAB Reservoir Simulation
 307 Toolbox, MRST² (Lie, 2019; Lie et al., 2012). A flowchart of the proposed approach is
 308 illustrated in Figure 8.

² <https://www.sintef.no/projectweb/mrst/>

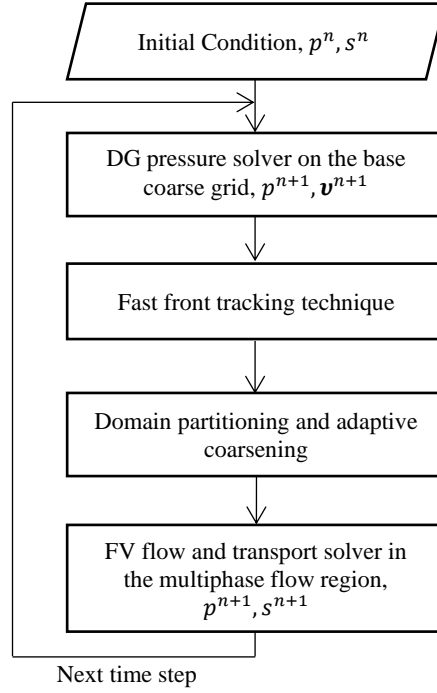


Figure 8 The flow chart of the proposed sequential approach for each time step

309 4.1 Complexity analysis

310 We use the number of unknowns solved at each step as the main indicator of the complexity of the
 311 solution approach so that this analysis does not depend on the actual implementation of the method. In
 312 the following, we detail the computational cost of each step separately.

313 (1) DG pressure solver on the base coarse grid: The number of unknowns at this step is equal to
 314 the number of coarse grid blocks times the number of local degrees of freedom which is equal
 315 to:

$$N = (P + 1)^2 \quad \text{Eq. 30}$$

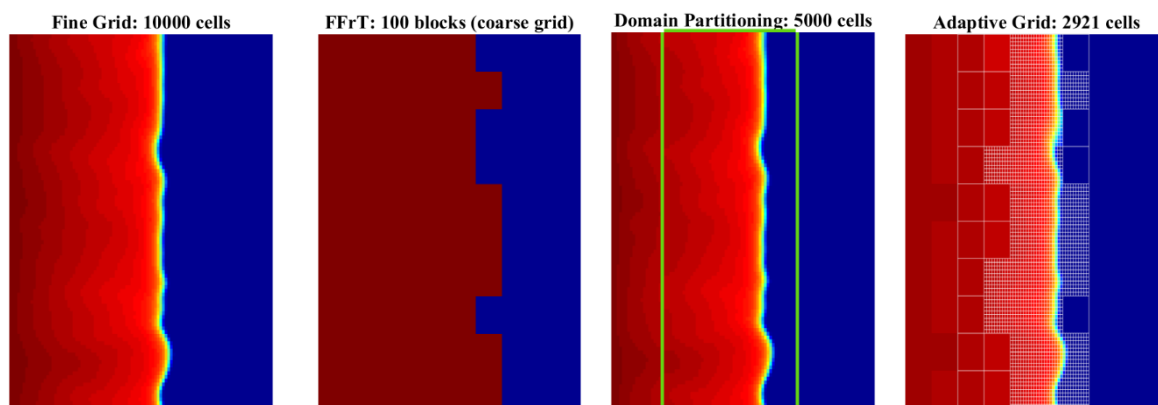
316 where P is the polynomial degree.

317 (2) The FFrT method: In this step, Eq. 28, is solved in parts of the domain where the front is present,
 318 and the number of unknowns here depends on the shape of the front. Thanks to these two
 319 important features, being local to cells where the front is present, and the explicitness of the
 320 equation to be solved, this step is computationally very fast.

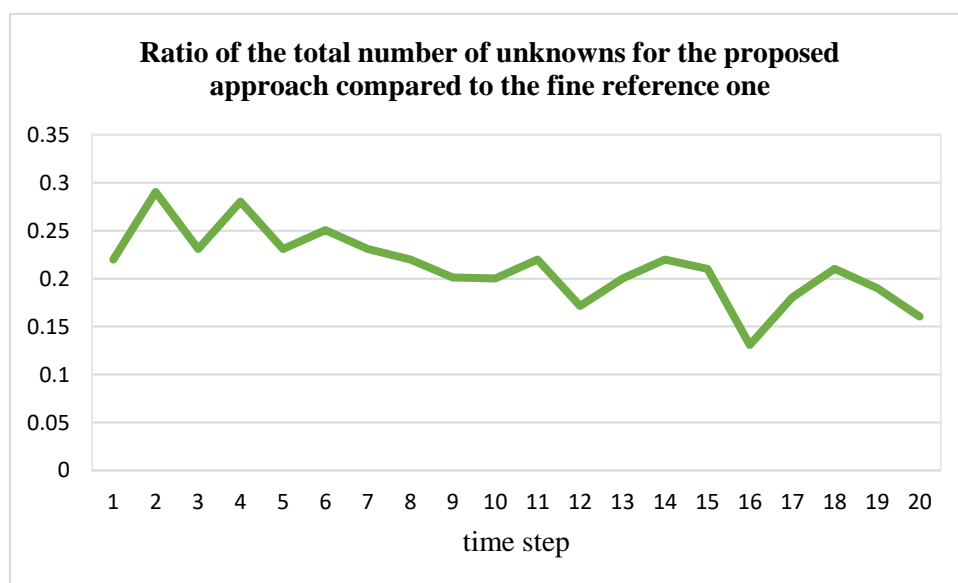
321 (3) Adaptive coarsening: In the agglomeration-based coarsening approaches, a coarse grid is
 322 generated from the agglomeration of the fine grid using a partition vector to relate coarse blocks
 323 to their underlying fine cells. Thanks to this preserved one-to-one mapping between fine and
 324 coarse grids, adding or removing local resolution is simple to carry out. Even if the cost of the
 325 grid adaptations is difficult to quantify precisely, in test cases we performed, this step was less
 326 than about five per cent of our overall computational time. It is also worth noting that this
 327 dynamic coarsening reduces the number of unknowns in the computationally demanding
 328 transport solver.

329 (4) FV pressure and transport solver in the refined areas: The number of unknowns in this step is
 330 twice the number of cells in the two-phase flow region.

331 As an illustrative example, we test the method with the random permeability field of Figure 3. Water
 332 and oil viscosities are equal to 1.0 and 0.2 cp, respectively. The reservoir is initially filled with oil. A
 333 coarse grid is generated by uniform agglomeration of the original grid with a coarsening ratio of 10.
 334 This simple example with sharp and smooth fronts shows how the proposed approach works. Figure 9
 335 shows the resulting saturation maps along with the fine solution taken as reference. The second plot,
 336 from left, shows a pseudo saturation map that indicates the presence or absence of water given by the
 337 FFrT method. The third plot shows the results of the domain partitioning scheme, where the saturation
 338 equation is solved using the FV method only in the identified two-phase region with a high-resolution
 339 grid. The green lines show the boundaries of the two-phase region. This domain is only defined from
 340 its west and east boundaries. The north and south boundaries coincide with the main boundaries of the
 341 reservoir. The east boundary is determined thanks to the pseudo saturation map. Since this criterion
 342 cannot be used to determine the west boundary of the domain, the saturations calculated from the
 343 previous time step are used instead. The last plot in this figure shows the results obtained with our
 344 procedure, where both domain partitioning and dynamic nonuniform coarsening are used. Figure 10
 345 shows the ratio of the total number of global degrees of freedom in the proposed approach compared to
 346 the reference solution for this example.



347 Figure 9 From left to right: the water saturation map on the base fine grid taken as the reference solution,
 348 the pseudo-saturation map on the coarse mesh based on FFrT method, the water saturation map for the
 349 domain partitioning scheme (without adaptive coarsening), and the results obtained with the proposed
 350 approach using domain partitioning and adaptive coarsening. The green line indicates the two-phase
 351 partition of the domain. The number of grid cells is shown above each plot.



352 Figure 10 The ratio of the total number of unknowns to be solved at each time step in the proposed
 353 solution approach compared to the fine reference solution.
 354

355 5 Numerical Examples

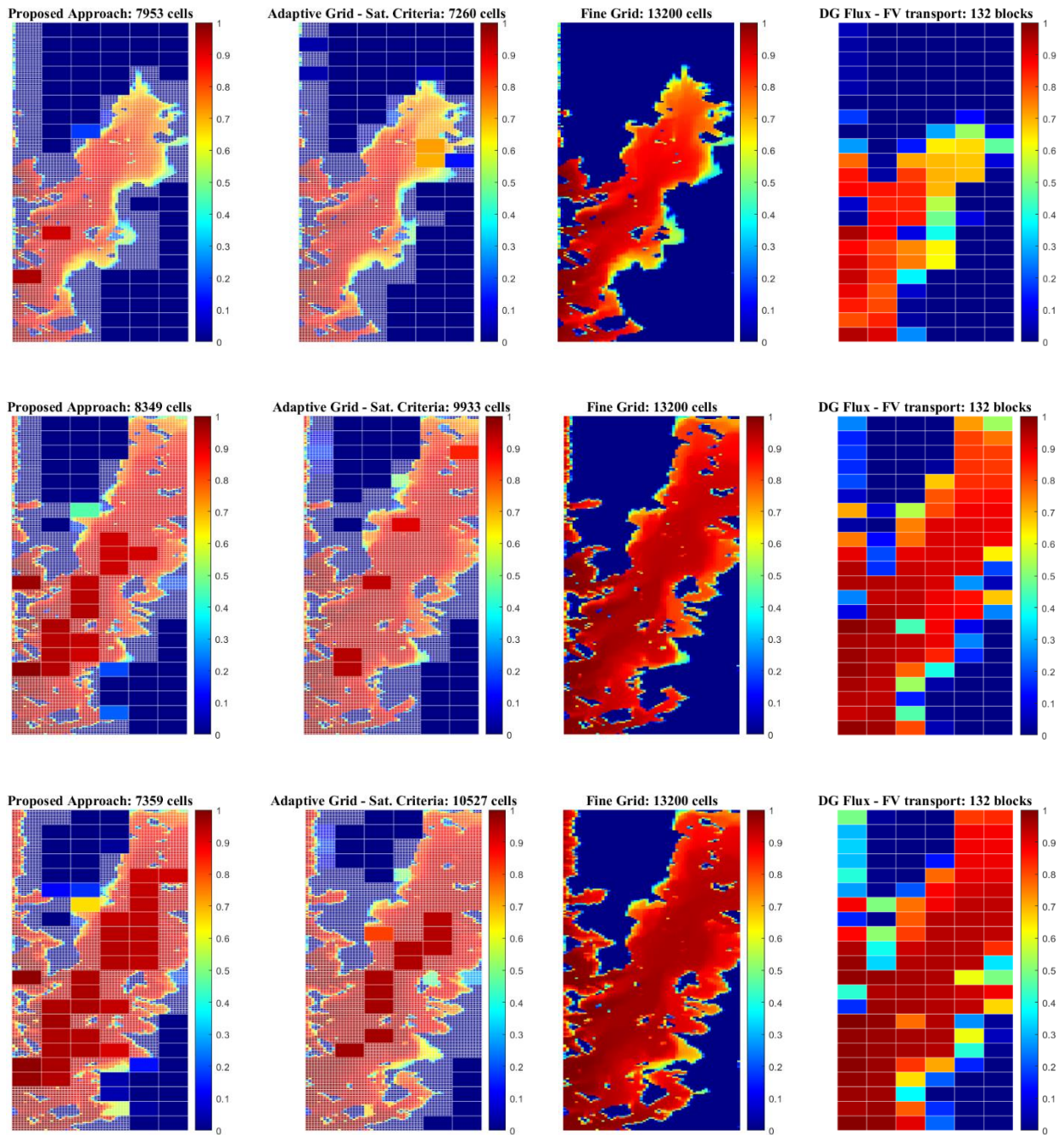
356 In the following examples, we test the approach in two-dimensional waterflood problems with
357 favourable and unfavourable viscosity ratios. The initial and boundary conditions remain the same as
358 in the previous examples. In the first two examples, the identified coarse blocks are replaced with their
359 underlying fine grid cells. In the third example, we present cases where the resolution of the identified
360 coarse blocks are increased to an intermediate level.

361 5.1 Example 1: Favourable displacement

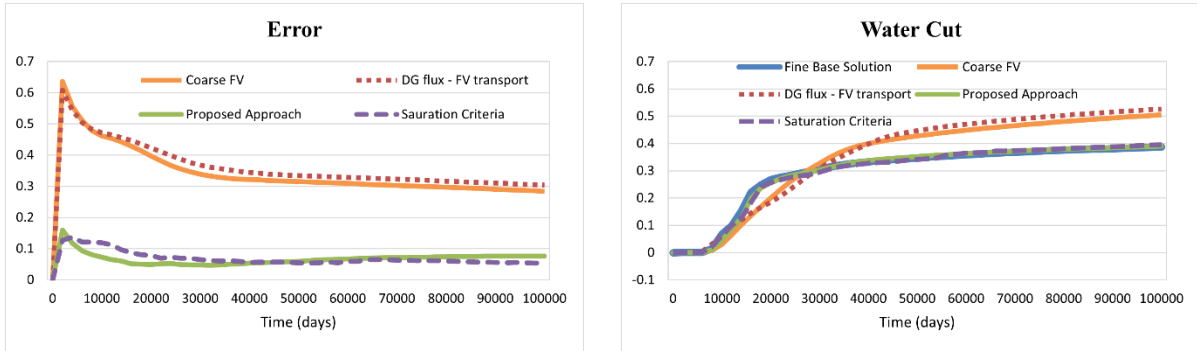
362 We test the proposed solution strategy in a favourable displacement process in a two-dimensional
363 Cartesian model with permeability values taken from layer 70 of the second SPE10 benchmark test.
364 Figure 11 shows the water saturation maps at three different time steps in this channelized reservoir.
365 The viscosity of water is 1.0 *cp* and the viscosity of oil is 0.2 *cp*. We can see that our proposed approach
366 captures the channelized flow detail which is challenging to get when a low-resolution grid is used over
367 the whole domain. In this figure, we also compare the proposed FFrT technique with a classic criterion,
368 where the saturation equation is solved on a coarse grid and the saturation change from the previous
369 time step is used to locally increase the mesh resolution in the blocks where this change exceeds a
370 defined tolerance. The FV solution on the original fine grid is also shown as a reference. To evaluate
371 the accuracy of DG conservative fluxes, the results of the finite volume transport solution using DG
372 computed total fluxes on the coarse grid are computed. We can see in this example that FFrT method
373 gives a better indication of where to add a higher spatial resolution. The proposed FFrT method solves
374 an explicit equation on some parts of the domain, while the saturation gradient criterion needs the
375 solution of the transport equation over the whole domain. To evaluate the accuracy of the proposed
376 approach, we use the L^1 relative error norm to compare the proposed solution strategy with the reference
377 fine solution, computed using the equation below:

$$err_h = \frac{\|s_h - s_{ref}\|_{L^1(\Omega)}}{\|s_{ref}\|_{L^1(\Omega)}} \quad \text{Eq. 31}$$

378 where s_{ref} is the reference solution where finite volume is used for solving both equations on the
379 original fine grid, and s_h represent the saturation in the chosen solution strategy. A comparison of the
380 errors and water cuts obtained with different schemes are shown in Figure 12. The proposed scheme
381 gives satisfactory results in terms of error, water breakthrough time, and the water cut. In these plots,
382 the results of the finite volume solution on the coarse grid are shown as well. In terms of error, the DG
383 conservative fluxes give results close to the finite volume coarse resolution. Figure 13 shows the ratio
384 of the total number of degrees of freedom for each time step for the proposed approach compared to the
385 base fine solution. The number of unknowns and subsequently the computational efficiency of the
386 approach depends on the extension of the multiphase flow area over the domain. It also depends on the
387 spatial resolution of the refined subdomains. In this example, water breakthrough happens early during
388 the simulation and the multiphase subdomains are replaced with their underlying fine grid cells.

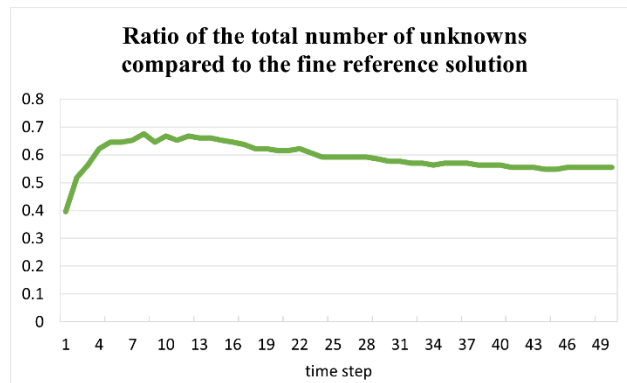


389 Figure 11 Saturation maps for the layer 70 of the SPE10 model computed by the proposed approach,
 390 the adaptive approach with saturation gradient criteria, FV on the base fine grid, and DG(flux)-
 391 FV(saturation) schemes, at three different time steps.



392

393 Figure 12 Errors and water cut as a function of time for the layer 70 of the SPE10 model in a favourable
 394 displacement process

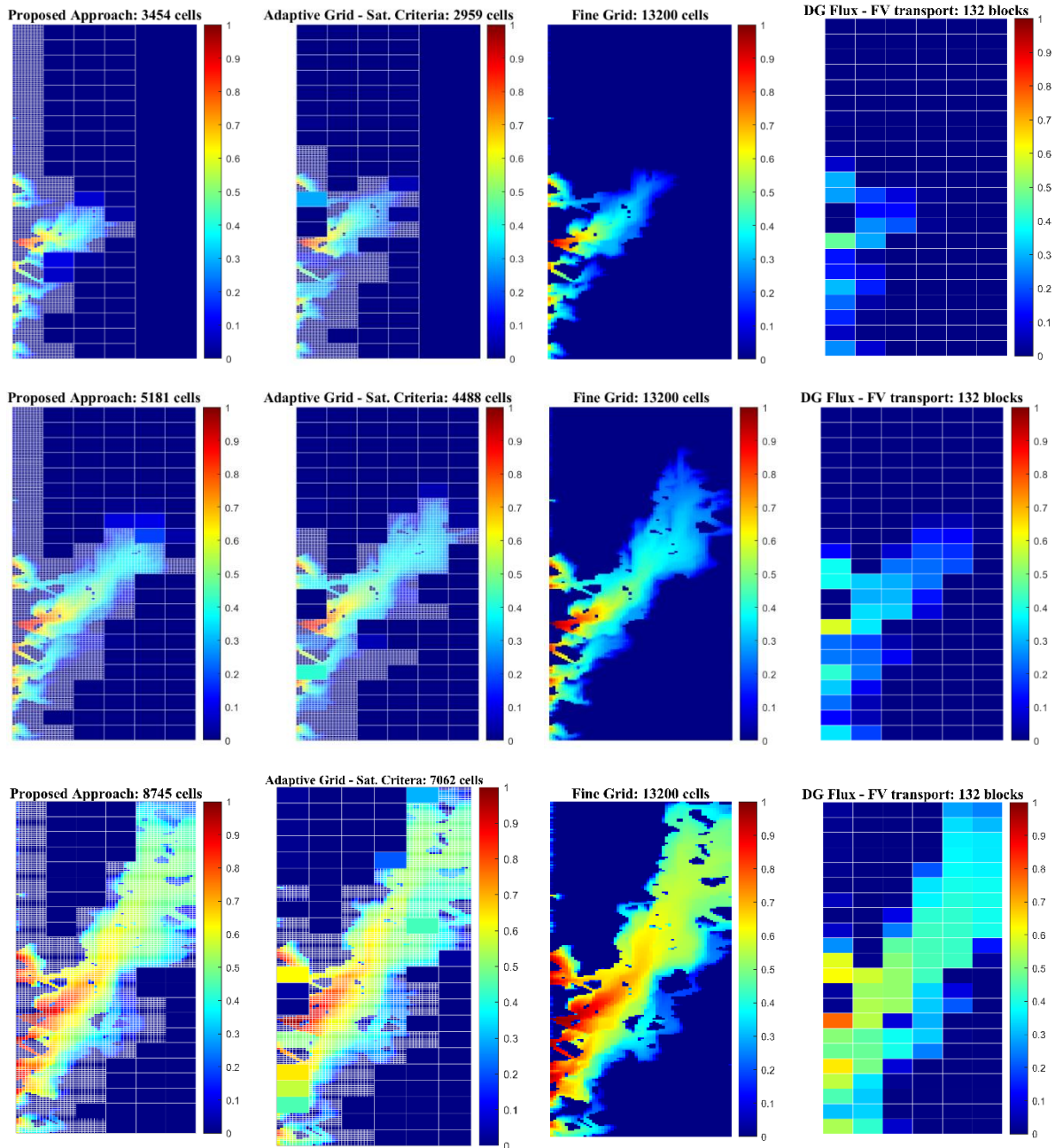


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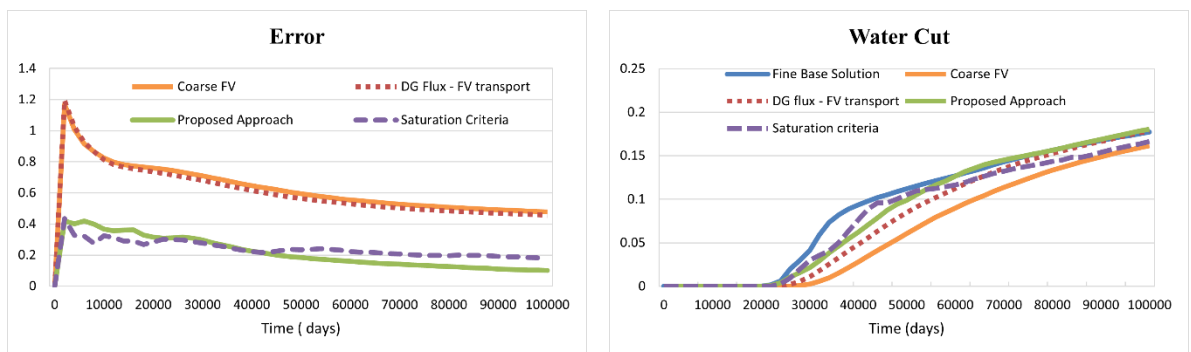
396 Figure 13 The ratio of the total number of unknowns for each time step in the proposed approach
 397 compared to the base fine solution

398 **5.2 Example 2: Unfavourable displacement**

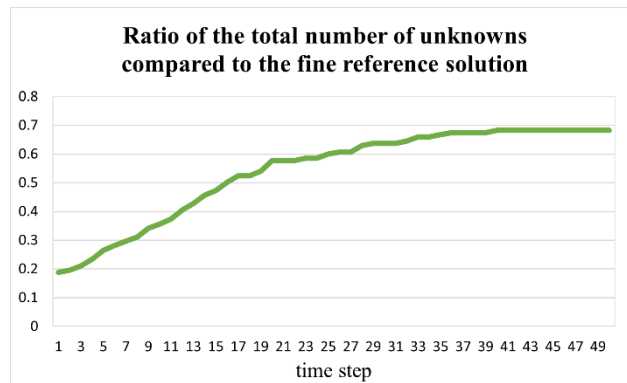
399 To test the proposed solution strategy in unfavourable viscosity ratios, we set the viscosities of water
 400 and oil to 0.5 and 5 *cp*, respectively, to get an unstable viscous flow at the front. All other conditions
 401 remain the same as in the previous case. Figure 14 shows the water saturation maps at three different
 402 time steps for the same layer. Our proposed approach successfully captures the diffusive flow pattern
 403 in the channel. Water saturation gradient works slightly better than the FFrT in the early time steps but
 404 tends to decline after the water breakthrough time. Figure 15 shows the errors and water cuts for
 405 different schemes. We notice that in this case, the coarse DG conservative fluxes give slightly better
 406 results than the coarse FV fluxes, in terms of the water cuts. The dynamic coarsening approaches with
 407 FFrT method and the classical criterion of saturation gradient give similar results in terms of relative
 408 error and water cut. However, the FFrT method is much faster than any transport solver. This figure
 409 also shows the relatively large errors in coarse-scale solutions. Figure 16 shows the total number of
 410 global degrees of freedom for the proposed approach relative to the reference fine solution. The
 411 proposed approach has an accuracy close to the reference solution while decreasing the number of
 412 global degrees of freedom.



413 Figure 14 Water saturation maps for the layer 70 of the SPE10 model, at three different time steps
 414 before (first row), at (second row), and after (third row) the water breakthrough for an unfavourable
 415 displacement case.



416 Figure 15 Errors and water cuts as a function of time for the layer 70 of the SPE10 model with an
 417 unfavourable viscosity ratio.

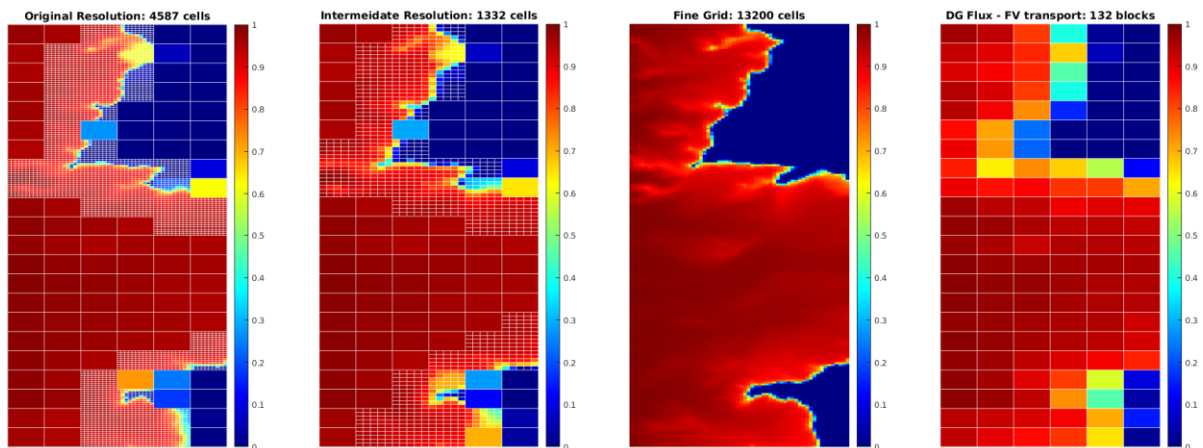


418
 419 Figure 16 The ratio of the total number of unknowns for each time step in the proposed approach
 420 compared to the base fine solution for the unfavourable displacement in layer 70 of SPE10 benchmark
 421 test

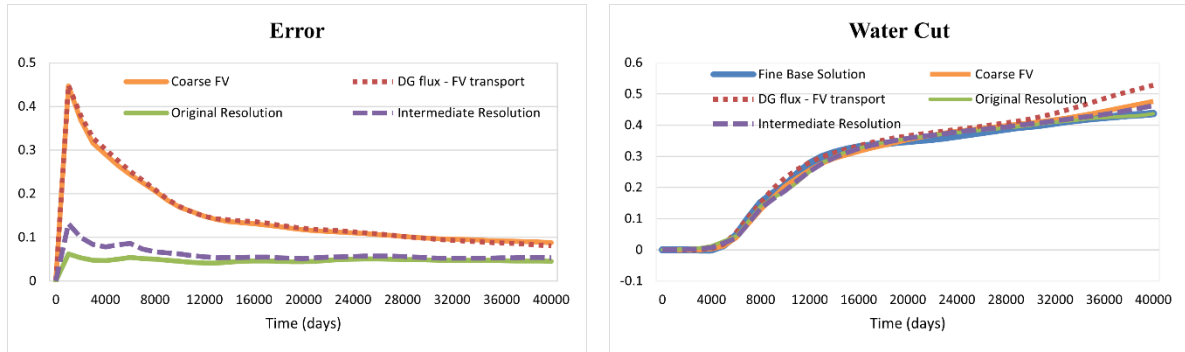
422

423 5.3 Example 3: Intermediate resolution

424 In the previous examples, we replaced the coarse blocks with their original fine resolution grid cells in
 425 the adaptive coarsening step. But we can also increase the spatial resolution in the indicated coarse grid
 426 blocks to any intermediate resolution, to decrease the number of degrees of freedom and ultimately
 427 improve the computational efficiency. This is especially more efficient in displacements with sharp
 428 saturation fronts. In Figure 17 you can see the results of the approach using an intermediate spatial
 429 resolution, where the identified coarse blocks are replaced with a finer resolution of ratio 2 relative to
 430 the original resolution. Here the permeability field is taken from layer 22 of the SPE10 second model
 431 and viscosities of water and oil are set to 1.0 and 0.2 cp, respectively. The errors and water cuts are
 432 represented in Figure 18.

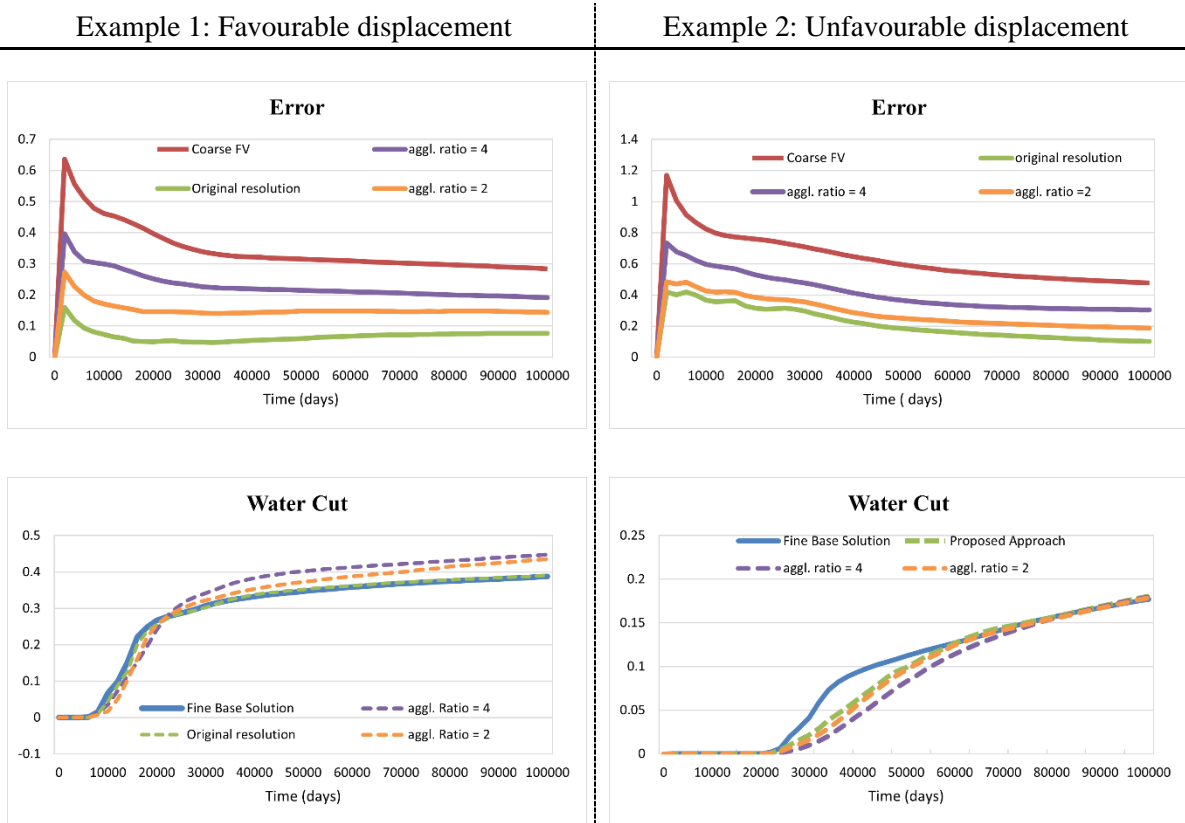


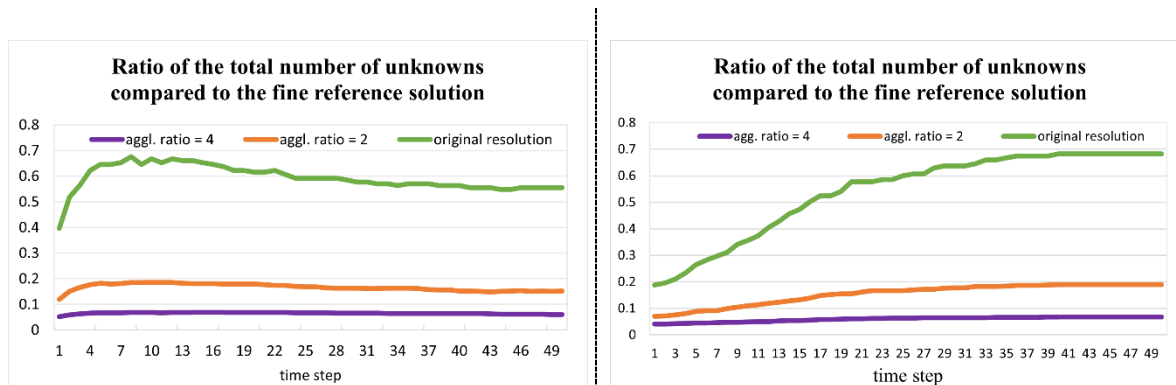
433
 434 Figure 17 Saturation maps for layer 22 of the SPE10 with a favourable viscosity ratio, for two different
 435 fine resolutions.



436 Figure 18 Errors and water cuts as a function of time for the layer 22 of the SPE10 model and two
 437 different intermediate resolutions.

438 We revisit the examples 5.1 and 5.2 and we increase the resolution of identified coarse blocks to
 439 different intermediate levels with agglomeration ratios of 2 and 4 relative to the base fine grid. Figure
 440 19 shows the error, the water cut, and the total number of unknowns for both examples. In this figure,
 441 the original resolution represents the solution where the coarse blocks are replaced with their original
 442 fine cells. This figure shows that replacing the coarse blocks with an intermediate resolution of the
 443 agglomeration ratio 2 gives suitable accuracy while decreasing the total degrees of freedom to a great
 444 extent.





445 Figure 19 The errors, the water cuts, and the ratio of the total number of unknowns compared to the fine
 446 reference solution for the previous examples with different intermediate resolution level.

447

448 6 Conclusions

449 In this paper, an original method combining dynamic non-uniform coarsening and a discontinuous
 450 Galerkin method was developed. The goal is to improve the efficiency of multiphase flows simulations
 451 in heterogeneous porous media without losing accuracy. The proposed fast front tracking method
 452 appears as being a promising method to combine different resolution strategies focusing on different
 453 flow areas. This method can be used on a coarse scale to identify the two-phase flow region with
 454 satisfying accuracy and a small computational cost of solving an explicit equation in some parts of the
 455 domain. This information can thus be used as input to dynamic mesh coarsening and adaptive use of
 456 DG and FV solvers for the pressure equation. This method has proved to be a powerful tool to predict
 457 the position of the front. Near the front, in the two-phase region, we have shown that a high-resolution
 458 grid used along with a finite volume discretization leads to stable solutions with improved accuracy.
 459 Far from the front, the DG method, used on a lower resolution grid, increases the accuracy of the total
 460 velocity.

461 The efficiency of the approach depends on the overall spreading of the multiphase region, and the level
 462 of resolution in the adaptive coarsening. We have shown that the resolution level can be adjusted
 463 depending on the required order of accuracy, the available computational cost, and the complexity of
 464 the problem.

465 However, the proposed solution strategy can be extended to a more general framework. The fast front
 466 tracking method can be improved to handle more complex flows like radial ones around wells in 3D
 467 domains. Other approaches could also be followed. For example, working at a coarse-scale using
 468 classical single-phase upscaling, coupled with up-scaled two-phase flow equations with an effective
 469 fractional flow function (Artus et al., 2004; Fayers et al., 1992; Sorbie et al., 1995) and a
 470 macrodispersion term modelling the subgrid disorder. The effective fractional flow accounts for the
 471 average local pressure saturation coupling. This could help to get a fast estimation of the front location
 472 and its typical thickness at the coarse scale directly. If necessary, a mesh refinement will then be set up
 473 in that area depending on a posteriori criterion quantifying the overall accuracy of the calculation
 474 (Gratien et al., 2016).

475

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